

# DSD variability: normalization and retrieval

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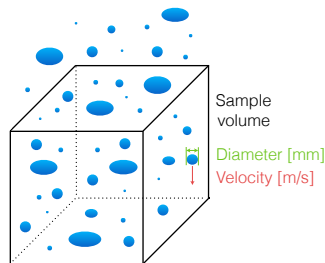
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# The raindrop size distribution (DSD)

- Concentration in air of raindrops with equivolume diameter in  $[D, D + dD)$ :

$$N(D) dD$$

- Weighted moments  $\rightarrow$  rainfall variables.
- Measured at point scale by disdrometers.
- DSD variable in space and time  $\rightarrow$  various normalizations.



# Double-moment normalization of the DSD

The  $n^{\text{th}}$  moment of the DSD is  $M_n = \int_0^{\infty} D^n N(D) dD$  [mm <sup>$n$</sup>  m<sup>-3</sup>]

With moment orders  $i$  &  $j$ , DSD can be written (**Lee et al. JAM 2004**):

$$N(D) = M_i^{(j+1)/(j-i)} M_j^{(i+1)/(i-j)} h(x),$$

where  $x$  is a normalized diameter:

$$x = D M_i^{1/(j-i)} M_j^{-1/(j-i)}.$$

**To estimate the DSD we need three ingredients:**  $h$ ,  $M_i$  and  $M_j$

- How variable is the double-moment normalized DSD  $h$  in space?
- Can the double-moment normalization be used for radar DSD retrieval?

# Instrument networks

Maps ©Thunderforest (CC BY-SA), map data ©OpenStreetMap (ODbL)

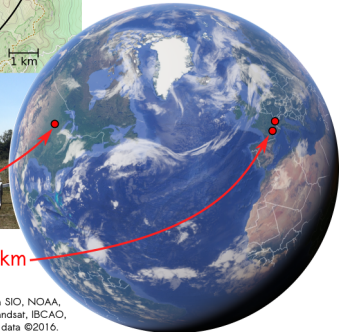


Horizontal



Vertical

> 7000 km



Instrument networks:

Ardèche, France

Payerne, Switzerland

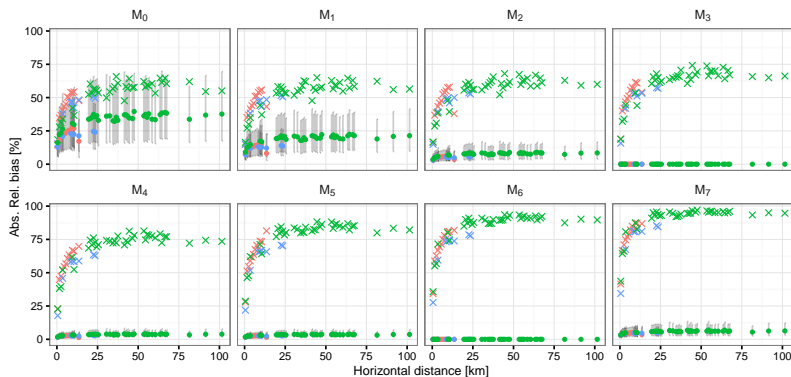
Iowa, United States

~ 85% stratiform rain.

Google. Imagery ©2016 Data SIO, NOAA,  
U.S. Navy, NGA, GEBCO, Landsat, IBCAO,  
U.S. Geological Survey, Map data ©2016.



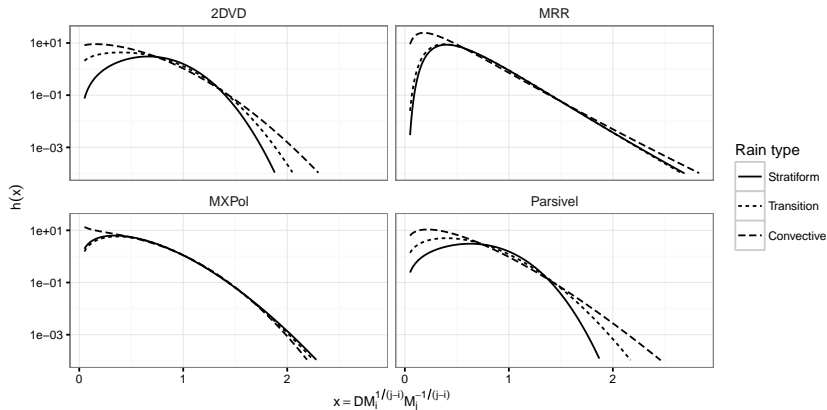
# Changes in the horizontal



Function ☐  $h(x)$  ☒  $DSD$  Network ☒ HyMeX ☒ Iowa ☒ Payerne  $i = 3, j = 6$

- normalized DSD “collapses” spatial variability.
- Similar for vertical variability (although larger remaining variability).

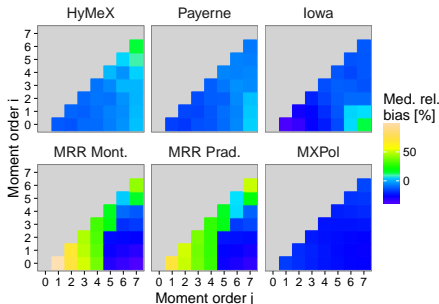
# One fitted model per instrument type



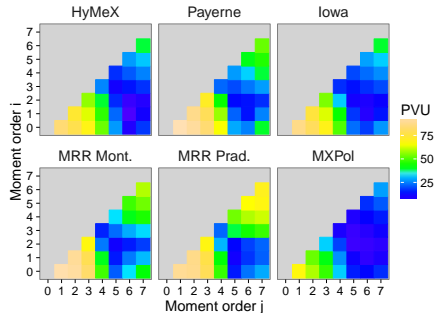
- Models trained on HyMeX (France) data only.
- Can they be used in other regions?

# Performance by region in stratiform rain

Median rel. bias



% variance unexplained



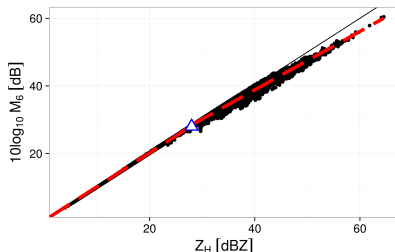
- Model trained in France can be applied to other regions.
  - Best performance with varied moment orders (e.g. 1 and 6).
  - With best combination,  $\geq 85\%$  of DSD variance can be explained.
- ⇒ For DSD retrieval,  $h$  can reasonably be assumed invariant in space.

# DSD retrieval

We have  $h$ , so we need two moments to retrieve the DSD.

## Moment 6

- Retrieved using  $Z_h$  power laws.
- Split into two regimes at 28 dBZ (X-band).



## Moment 3

- Z-weighted mean drop axis ratio  $\hat{r}_m$  estimated using  $Z_{DR}$ .\*
- $M_3$  retrieved using  $\hat{r}_m$  and  $K_{dp}$ .
- Coefficients per raindrop axis ratio function.

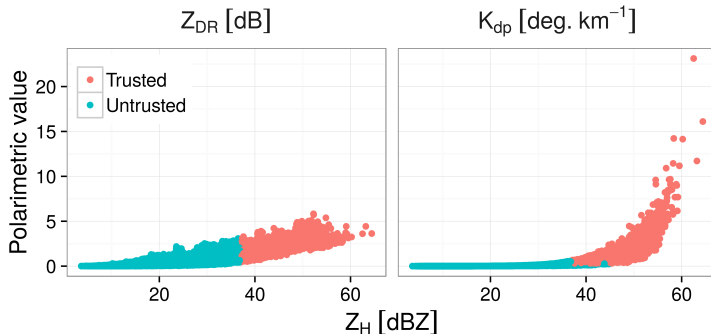
## Parameterization

- 60% of DSDs from HyMeX, Payerne, and Iowa.
- Simulated variables at X-band.
- Temps. of 5, 10, and 15°C.

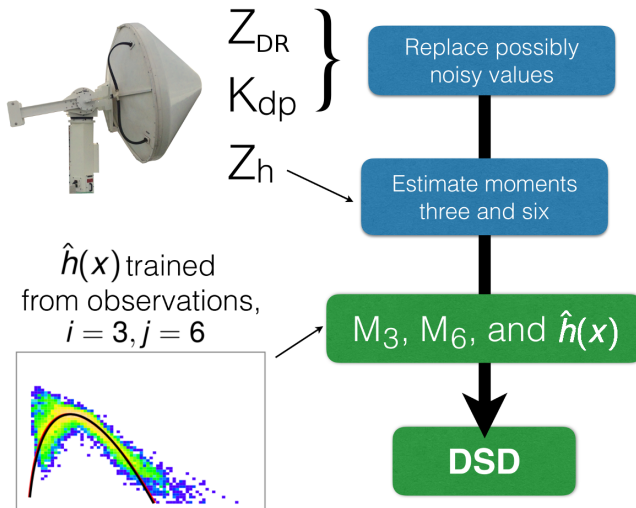
\* See Jameson JAS 1983, Kalogiros et al, IEEE TGR 2013.

# Dealing with radar noise

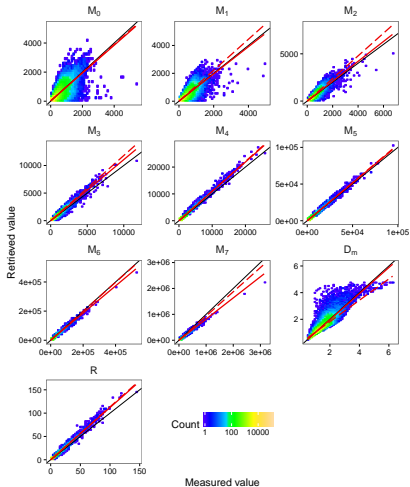
- $Z_{DR}$  and  $K_{dp}$  can be noisy.
- Threshold on  $Z_H < 37$  dBZ,  $K_{dp} < 0.3$  ° km<sup>-1</sup>,  $Z_{DR} < 0.2$  dB.
- In this case,  $Z_{DR}$  and  $K_{dp}$  predicted from  $Z_H$ .



# DSD-retrieval method



# Results with simulations

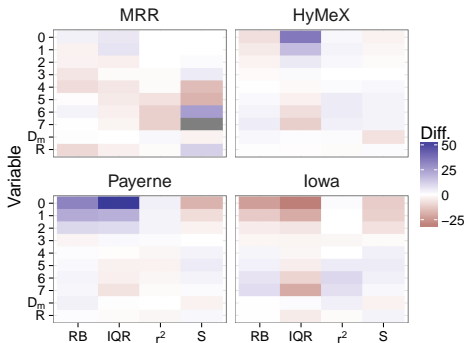


HyMeX results, Beard axis ratios.

- Simulated radar variables from disdrometer measurements.
- Compared against SCOP-ME (Anagnostou et al. AR 2009, Anagnostou et al. JH 2010, Kalogiros et al. IEEE TGRS 2013).
- Four different raindrop axis ratios (Beard 1987, Andsager 1999, Brandes 2002, Thurai 2007).
- Performances similar, on average DM has slightly lower bias for moments two to seven,  $D_m$ ,  $R$ .

# Results with real radar data

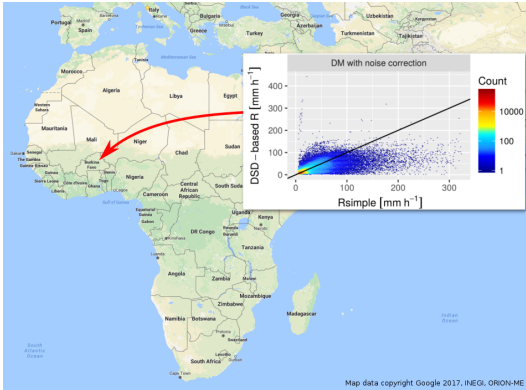
- Compared against Parsivels on ground in HyMeX, Payerne, and Iowa.
- In HyMeX, also compared against DSDs estimated aloft by an MRR (DM model trained for MRR data).
- Similar results for 2 algorithms; DM slightly better on MRR and Iowa data.



Reds show performance better with DM.



# Tests in Africa



- X-band radar data from Burkina Faso.
- Provided by M. Gosset and M. Kacou (IRD, Toulouse, France).
- $h(x)$  retrained using local disdrometers.
- Promising results: DSD-based  $R$  matches closely to Z-R-based  $R$ .

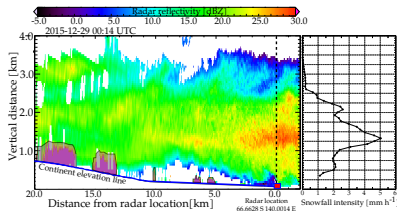
# Conclusions

- Relative invariance of the double-moment normalized DSD  $h(x)$ .
- A new DSD-retrieval technique based on the double-moment normalization approach.
- Flexible, since there is no prescribed form of  $h(x)$ .
- Future work: remaining variability in  $h(x)$ , other radar wavelengths, other rain climatologies.

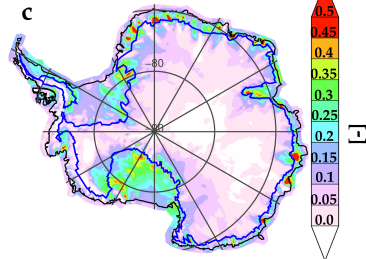
Raupach & Berne, *Invariance of the double-moment normalized raindrop size distribution through 3D spatial displacement in stratiform rain*, JAMC, 2017, 10.1175/JAMC-D-16-0316.1

Raupach & Berne, *Retrieval of the raindrop size distribution from polarimetric radar data using double-moment normalization*, AMT, 10, 2573-2594, 2017, amt-10-2573-2017

# Precipitation and katabatic winds in Antarctica



Ratio of sublimated snowfall



Significant low-level sublimation of snowfall  
(Grazioli et al., PNAS, 2017)

Poster initially planned yesterday, but now displayed on board #207 (thanks Walt!)

# Thank you



Rainfall over mountains in Ardèche during HyMeX 2012 SOP

## Moment 3 retrieval (1)

1. Retrieve radar-weighted mean raindrop axis ratio  $r_m$  [-] from  $Z_{\text{DR}}$ , using polynomial fit:

$$\hat{r}_m = \sum_{i=0}^5 c_i Z_{\text{DR}}^i.$$

2. Use relationship between LWC and  $M_3$  to derive (338.4 for X-band):

$$M_3 = \frac{338.4}{\hat{C}} \frac{K_{\text{dp}}}{(1 - \hat{r}_m)},$$

3. Polynomial fit and representative (mean) value of  $\hat{C}$  are parameterized per axis ratio function.

## Moment 3 retrieval (2)

Liquid water content  $W$  [ $\text{g m}^{-3}$ ] (with water density  $\rho_w$   $\text{g cm}^{-3}$ ):

$$W = \frac{\pi}{6} 10^{-3} \rho_w M_3,$$

and  $K_{\text{dp}}$  can be written (Jameson JAS 1985) :

$$K_{\text{dp}} = \left( \frac{180}{\lambda} \right) 10^{-1} C W (1 - r_m),$$

with  $C \sim 3.75$  (Bringi and Chandrasekar, 2001).